

# New estimation models for determining the $Q_{347}$

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EGU Session: HS2.4.3

## Introduction

In Switzerland, low flows are described by the five percent quantile denoted by  $Q_{347}$ . This threshold value not only has consequences for the planning, but also necessitates authorities to adjust the operation of pertinent infrastructure to mitigate ecological impacts on watercourses. In accordance with Swiss federal law (1991), determination of the  $Q_{347}$  must be done using the duration curve of a discharge time series spanning at least a ten-year period. Typically, said time series are not available for smaller catchments necessitating the estimation of the threshold value  $Q_{347}$ . In Switzerland, the use of multiple linear regression has been established to estimate the area-specific discharge  $q_{347}$  (Aschwanden, Kan 1999). However, this regionalization method is associated with significant uncertainties for estimated values (Naef et al. 2015).

## Study area and data processing

The primary objective of these investigations is to estimate the  $Q_{347}$  value for 383 ungauged catchments in the Canton of Solothurn, each covering an area less than 100 km<sup>2</sup>. Daily discharge, precipitation and temperature timeseries ranging from 1990 to 2020 were collected from 56 gauged catchments smaller than 500 km<sup>2</sup> surrounding the target area. 30 «static» parameters describing geometry, topography, geology, land use, and drainage along with nine «climatic» parameters describing temperatures, precipitation distributions, and potential evapotranspiration were defined and computed to characterize gauged and ungauged catchments.

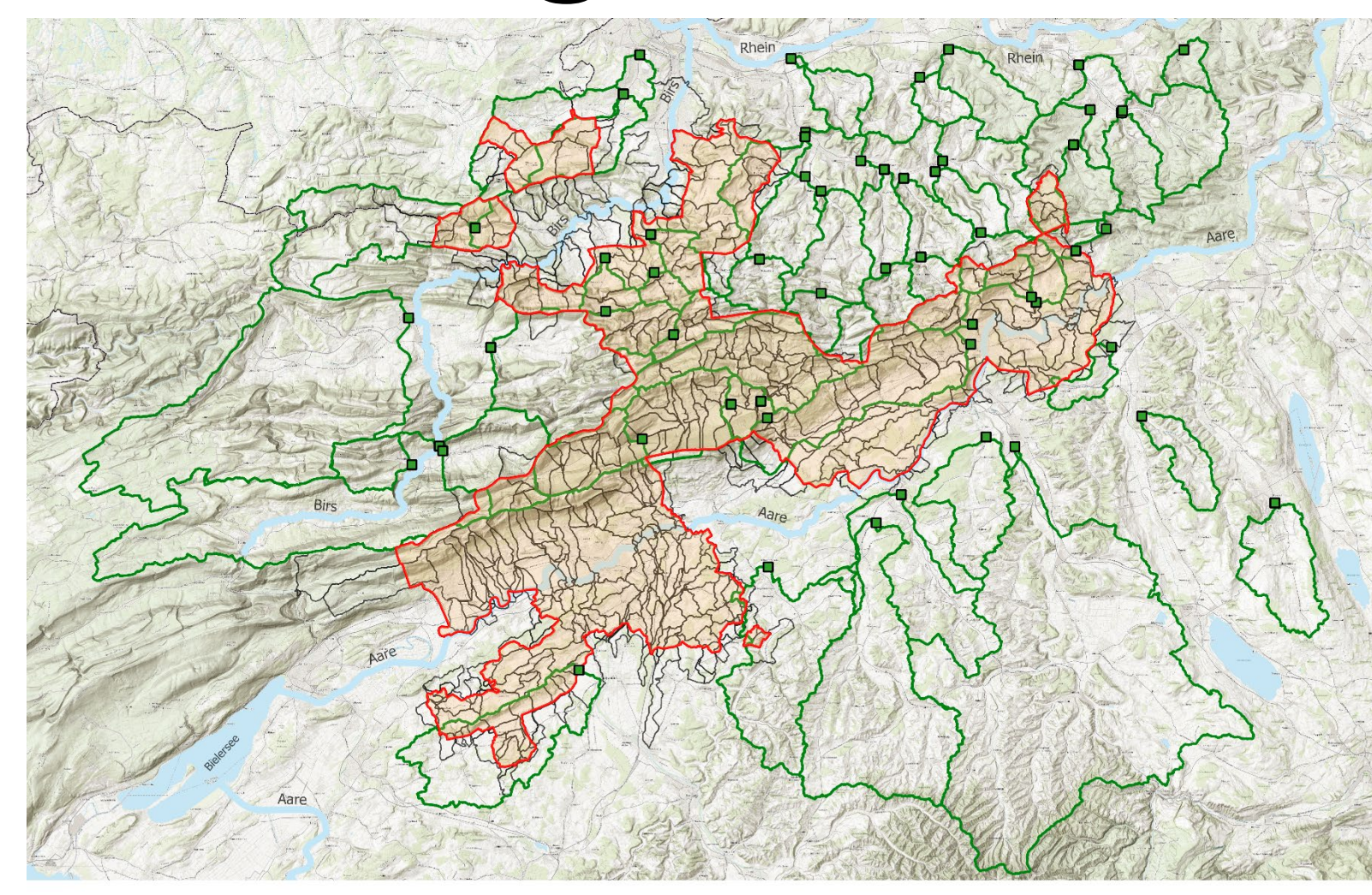


Fig. 1 Study area and discharge stations with gauged and ungauged catchments

## Previous investigations

### a) Variability of $Q_{347}$ over the study period and by decades

By definition, an average of 5 percent of all days must have  $Q \leq Q_{347}$ . Figure 2 shows the percentage of  $Q_{347}$ -days per year with a moving average of 5 years and illustrates a non-uniform distribution with an ascending trend in the frequency of low-flow events.

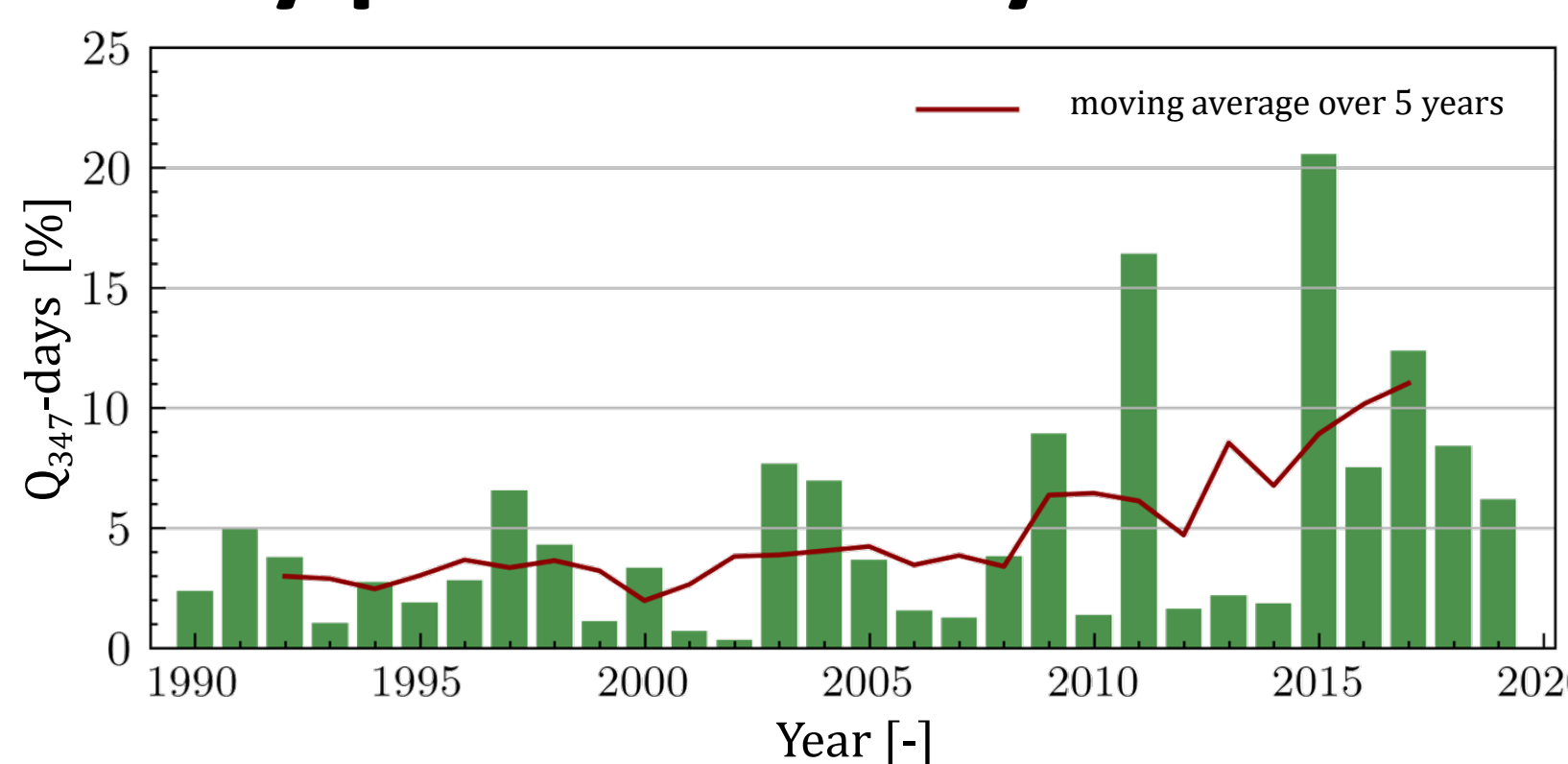


Fig. 2 Percentage distribution of  $Q_{347}$  days and 5 year moving average from 1990 to 2019.

Figure 3 illustrates the relative change in  $Q_{347}$  if the value is computed for each decade within the study period: 1990–1999, 2000–2009, and 2010–2019. Comparing the first two decades,  $Q_{347}$  is decreased on average by 7.2%. This negative trend is further intensified for the second two decades between 2000 and 2019: for 89 % of the gauged catchment  $Q_{347}$  decreased with an average decrease of -13.8%.

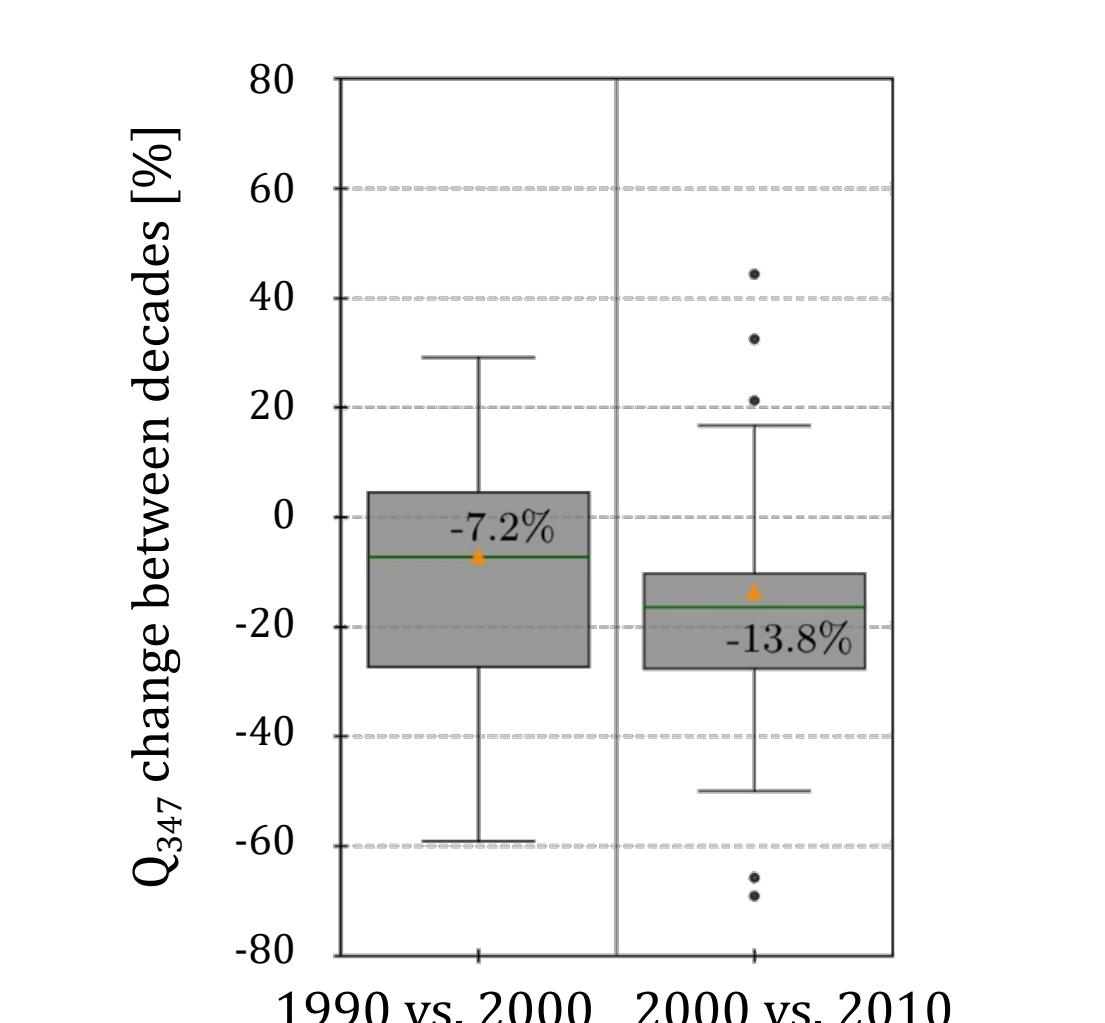


Fig. 3  $Q_{347}$  change between decades

### b) Short discharge time series as source of uncertainty

Figure 4 demonstrates the relative errors resulting from recalculated  $Q_{347}$  values after excluding data from randomly selected years within long-term time series. Short time series can be identified as significant source of uncertainty in estimating  $Q_{347}$ .

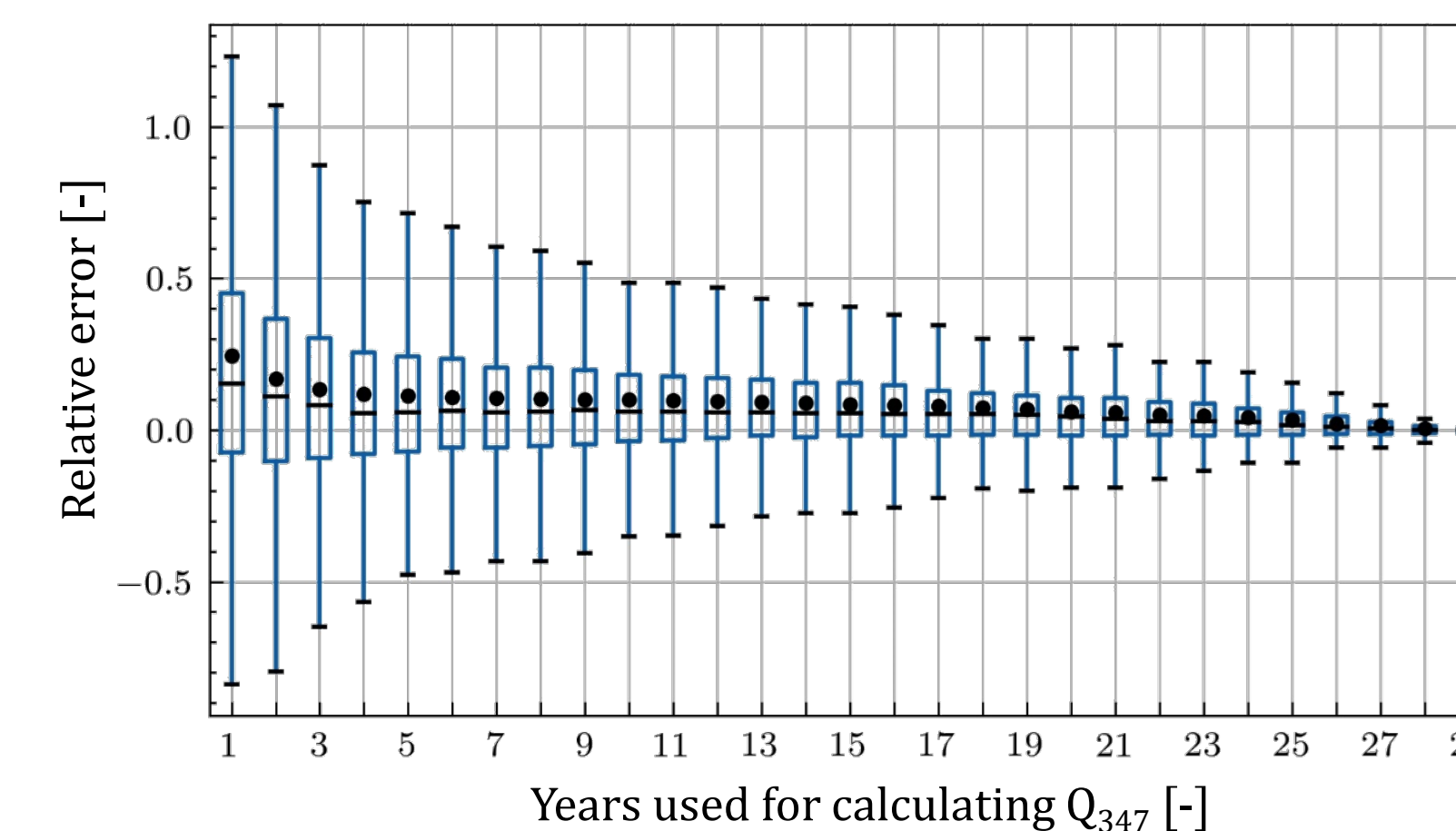


Fig. 4 Relative error in  $Q_{347}$  as a function of used data

## Regression models for estimating $Q_{347}$

### a) Regression models

Three regression models (M1 – M3), in combination with two methods (API, HEZG) for adjusting shortened time series (plus no adjustment), are proposed and compared with two reference models (BAFU and 123).

**M1 «Strong linear regression»:** More robust linear regression estimates can be achieved by simulating the linear regression, considering the variability of the observations (Pavia, 2022). The following steps are taken:

1. Fit linear model with 7 «best» parameters from «backward elimination»
2. Calculate residuals and take n of them  $r_i = y_i - \hat{y}_i$
3. Form Ensemble  $\{\hat{y}_1, \hat{y}_2, \dots, \hat{y}_n\}$ ,  $\hat{y}_i = \beta_0 + \beta_1 \cdot x + r_i$
4. Aggregate ensemble by mean  $\hat{y} = \frac{1}{n} \sum_{i=0}^n \hat{y}_i$

**M2 «Adopt model from gauged downstream catchment»:** A linear regression model is fitted for each discharge station to estimate upstream ungauged catchments. Therefore, the station is omitted to find the seven «best» parameters per station applying «backward elimination» with minimization of the leave-one-out cross-validation error. M2 consists of 56 models, each associated with one discharge station plus one model that covers the entire study area to estimate ungauged catchments not having a gauged downstream station.

**M3 «Clustering of catchments»:** Similar catchments are identified based on their characteristics using the Random Forest method (Tyrallis, 2019). Catchments are then divided into three clusters with a regression model considering the «best» parameters being fitted for each of them. Ungauged catchments are assigned to one of the three clusters and estimated with the corresponding model.

### b) Temporal adjustment of short discharge time series

**API:** The method (Ridolfi et al., 2020) extends the flow duration curve for missing years by using the corresponding exceedance probability of the Antecedent Precipitation Index (API).

**HEZG:** The closest downstream discharge station (HEZG) is used to correct the  $Q_{347}$  (Laaha & Blöschl, 2007) of the catchment with a short time series using (1).

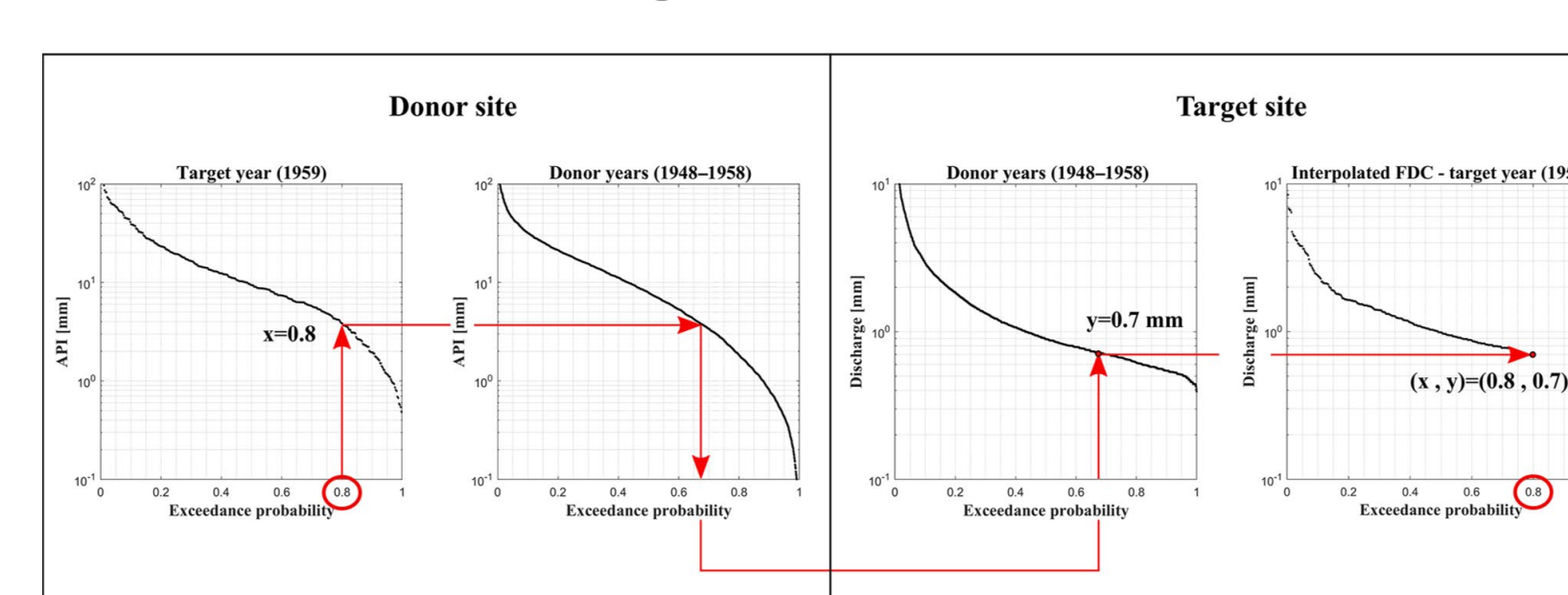


Fig. 5 Systematics of adjustment using API by the example of the 80th percentile (Ridolfi et al., 2020)

$$(1) \quad QT_{pred} = QT_O \left( \frac{QD}{QD_O} \right)$$

where:  $QT_{pred}$ :  $Q_{347}$  for target catchment in target periode  
 $QT_O$ :  $Q_{347}$  for target catchment in overlapping periode  
 $QD_O$ :  $Q_{347}$  for donor catchment in overlapping periode  
 $QD$ :  $Q_{347}$  for donor catchment in target periode

### c) Validation and regionalization

The proposed models were validated against the 56 gauged catchments and evaluated using three metrics: the linear correlation coefficient (Cor.), the mean error (ME), and the mean absolute percentage error (MAPE).

Model	New models									Reference models	
	M1			M2			M3			BAFU	1-2-3
Adjustment	No	API	HEZG	No	API	HEZG	No	API	HEZG		
<b>Metrics:</b>											
<b>Cor.</b> [-]	<b>0.77</b>	<b>0.73</b>	<b>-0.24</b>	<b>0.61</b>	<b>0.65</b>	<b>-0.29</b>	<b>0.46</b>	<b>0.37</b>	<b>-0.13</b>	0.16	-0.27
<b>ME</b> [L/s·km <sup>2</sup> ]	<b>0.03</b>	<b>0.08</b>	<b>12.5</b>	<b>-0.06</b>	<b>-0.02</b>	<b>12.52</b>	<b>-0.03</b>	<b>0.3</b>	<b>5.83</b>	-0.23	-0.75
<b>MAPE</b> [%]	<b>45</b>	<b>47</b>	<b>858</b>	<b>59</b>	<b>58</b>	<b>1045</b>	<b>55</b>	<b>62</b>	<b>416</b>	109	58

Table 1 Results of the validation on the gauged catchments (best three per metric are in bold)

For 383 ungauged catchments in the study area, the  $Q_{347}$  values were estimated using all three models in combination with the adjustment methods. The spatial distribution of the estimated  $Q_{347}$  values is shown in figure 6.

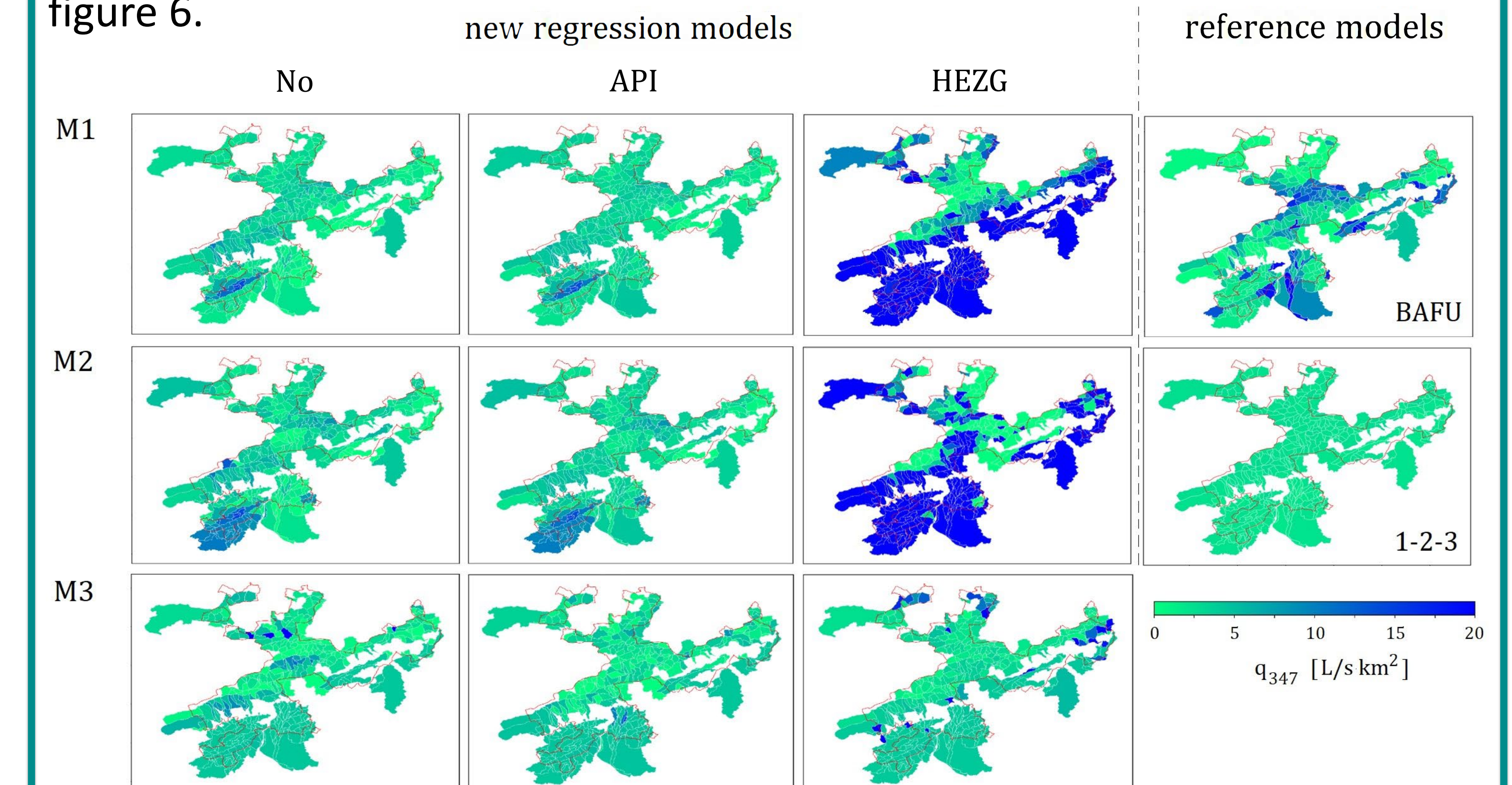


Fig. 6 Spatial representation of estimated  $Q_{347}$  values for 389 ungauged catchments applying nine proposed models and two reference models.

## Conclusions

- The frequency of low flow events below  $Q_{347}$  increased while the 10-year  $Q_{347}$  value of said catchments decreased over the last 30 years
- M1, M2, and M3 in combination with none or API adjustment for short time series improve the estimation of  $Q_{347}$
- Temporal adjustment using HEZG leads to large overestimations

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## Acknowledgments

- Amt für Umwelt des Kantons Solothurn.
- Data suppliers: Bundesamt für Meteorologie und Klimatologie MeteSchweiz, Bundesamt für Umwelt BAFU, Kanton Basel-Landschaft, Tiefbauamt, Hauptabteilung Wasserbau, Kanton Aargau Departement Bau, Verkehr und Umwelt, Kanton Solothurn, Amt für Umwelt, Bundesamt für Landestopografie swisstopo